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I e ed a e Wa e) (a e a e εNd = −8.2 0.3), LCDW (L e C c 1 R Dee Wa e) (a e a e εNd = −8.3 0.3),
NPDW (N z Pac c Dee Wa e) (a e a e εNd = −5.9 0.3), a d z e e a f NADW (N z A R c Dee Wa e)
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The REE can be divided into the Light REE (LREE) fraction, La/S, the Middle REE (MREE) fraction, HREE, and the Heavy REE (HREE) fraction. The REE patterns are similar to those of the corresponding REE minerals (e.g., Nya, 2001). REE concentrations are generally higher in the aluminosilicate minerals than in the feldspar minerals. The REE patterns are similar to those of the corresponding REE minerals (e.g., Díaz Redondo et al., 1990). The LREE fraction is enriched in La/HREE, and the HREE fraction is enriched in Nd/HREE. The REE patterns are similar to those of the corresponding REE minerals (e.g., Diaz Redondo et al., 1990).

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(Carvalho, 1972). However, the model of UCDW

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Tab 2

	De	La	Ce	P	Nd	Nd(ID)	S	E	Gd	Tb	D	H	E	T	Yb	L	HREE/ Ce LREE (E / Nd)	Ce/ LaR (Ce/Ce* = 2[Ce]/([La]+[P]))	
	()	(¹)	(¹)	(¹)	(¹)	(¹)	(¹)	(¹)	(¹)	(¹)	(¹)	(¹)	(¹)	(¹)	(¹)	(¹)	(¹)		
S a																			
S . 22(SO213-22-2) 39°12'S, 79°55'W	300	115	16.9	8.25	2.32	10.7	11.2	1.87	0.45	2.79	0.46	3.43	1.00	3.27	0.54	3.92	0.56	0.29	1.29
4144 de λ	650	107	15.6	5.84	2.36	10.6	10.4	1.71	0.45	2.36	0.43	3.86	1.14	4.02	0.59	4.24	0.71	0.39	0.65
	1500	148	22.3	6.78	2.87	12.7	12.9	2.16	0.61	3.23	0.55	4.84	1.44	5.36	0.86	5.98	1.14	0.42	0.54
	2600	192	28.6	7.15	3.55	14.3	16.1	2.39	0.79	4.20	0.62	5.77	1.76	6.00	1.08	6.87	1.29	0.37	0.45
	4142	178	33.8	7.74	4.68	21.0	21.0	3.70	0.92	4.94	0.85	6.85	2.00	6.85	1.06	8.02	1.36	0.33	0.40
S . 9(SO213-09-2) 37°41'S, 95°28'W	750	110	17.5	7.97	2.60	11.5	11.4	2.05	0.56	2.62	0.48	3.91	1.11	4.03	0.65	4.32	0.80	0.35	0.79
3771 de λ	1500	148	24.1	9.89	3.06	13.1	14.0	2.14	0.61	3.27	0.54	4.54	1.33	5.02	0.84	5.30	1.02	0.36	0.73
	2200	163	31.1	17.5	4.48	18.3	18.9	3.05	0.73	4.37	0.72	5.71	1.68	6.24	0.97	6.85	1.30	0.33	0.98

f æe d. d aRREE (Ha à e èl aR, 2012) de àcaR
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(2012).

2.2.3. Determination of nutrient concentrations

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3. RESULTS

1 Rca e c | cd | a | each ae | a | a | be ee | 1500 a d
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de λ . Be ed a - a eR 2000 ad 3000 de λ , λ e
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 a d e a laRPac cde ed a e a λ de λ a e,
 λ e e Nd c ce a >30 1 R a e e ec ed (Pe -
 a a d Jac b e, 1988; A a a a s AR 2004, 2009). T λ e
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3.4. Nd isotope compositions

T λ e Nd ec f λ e a 1 R a ef -
 -10.3 0.3 -5.4 0.4 ε_{Nd} (F . 5 TabR 1). T λ e -
 ed a e a e, λ e d a f AAIW ad ec
 SAMW (300 1500 de λ) λ a a e f ε_{Nd}

The ac f æ a a e a e (LCDW a d

-9.0 ± 0.3 (M C-78) and -10.3 ± 0.3 (M C-79) ϵ_{Nd} values were determined by the Nd isotope dilution mass spectrometry (ID-MS) method. The ϵ_{Nd} value of 1.3 was used for the M C-79 sample. The Nd isotope fractionation factor (ϵ_{Nd}) is defined as:

$$\epsilon_{\text{Nd}} = \frac{\text{Nd}_{\text{sample}}}{\text{Nd}_{\text{standard}}} - 1$$

where $\text{Nd}_{\text{sample}}$ and $\text{Nd}_{\text{standard}}$ are the measured Nd isotope ratios of the sample and standard, respectively. The Nd isotope ratio of the standard is taken as 0.51263.

(see Table 1 and Fig. 11e,f and 2b,c).

Figure 6 shows the Nd isotope composition of M C-79 ($\epsilon_{\text{Nd}} = -10.25$, $[Nd] = 21.8$ ppm), LCDW, and NADW, and the Nd isotope fractionation factor (ϵ_{Nd}) between LCDW and NADW at 3100 m, 179°E, 45°S. M C-79 is enriched in Nd relative to LCDW and NADW.

4.4. Sediment-bottom water interactions

4.4.1. Release of REEs from the sediments of the Southeast Pacific Basin

The release of REEs from the sediments of the Southeast Pacific Basin is controlled by the dissolution of authigenic minerals such as smectite and illite. The dissolution of these minerals is driven by the presence of organic matter and the reduction of oxygen. The dissolution of smectite and illite releases REEs into the bottom water, which can then be transported by the ocean currents. The dissolution of smectite and illite is also influenced by the presence of organic matter and the reduction of oxygen. The dissolution of smectite and illite releases REEs into the bottom water, which can then be transported by the ocean currents. The dissolution of smectite and illite is also influenced by the presence of organic matter and the reduction of oxygen.

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